

PROJECT

Numerical models for geothermal energy exploitation in

volcanic regions

Instructor Armando Coco Students Mattu Riccardo Marini Fabio

Catania

2024

Summary

- Physical problem
- Mathematical model
- Geothermal values
- Discretization
- Numerical results
- Conclusions and future works
- References

Physical problem





Figure: Geothermal power station

June 28, 2024 3/18



Equations

$$\begin{aligned} \frac{\partial T}{\partial t} + \mathbf{u} \cdot \nabla T - \alpha \Delta T &= 0 \quad (convection-diffusion \ equation) \\ \mathbf{u} &= -\frac{k}{\mu} \nabla p \qquad \qquad Darcy \ Law \\ div(\mathbf{u}) &= 0 \qquad \qquad Conservation \ of \ mass \end{aligned}$$

Mathematical model: Boundary condition





ECMI Mattu Riccardo Marini Fabio

June 28, 2024 5/18

Mathematical model: Temperature at t = 0



Figure: Plot of the temperature at the initial time

ECMI Mattu Riccardo Marini Fabio

June 28, 2024 6/18

E



coefficient	α	u	μ
physical meaning	diffusivity	permeability	viscosity
units	$m^2 \cdot s^{-1}$	m^2	$Pa \cdot s^{-1}$
l scenario	$1.43 \cdot 10^{-7}$	10^{-14}	$8.9 \cdot 10^{-4}$
II scenario	$2.3 \cdot 10^{-5}$	10^{-14}	$2.2 \cdot 10^{-4}$
III scenario	$0.5 \cdot 10^{-5}$	10^{-14}	$0.6 \cdot 10^{-3}$

- Diffusivity: this constant indicates how significant the diffusion is. The greater the alpha, the more the heat spreads.
- Permeability: is the ability of underground rocks to allow fluid to pass through.
- Viscosity: in a fluid are defined as those resulting from the relative velocity of different fluid particles.



We implemented two types of discretization for time dependance:

Equations (Forward Euler) $\frac{T^{n+1} - T^n}{\tau} + \mathbf{u} \nabla T^n - \alpha \Delta T^n = 0$

Equations (Backward Euler)

$$\frac{T^{n+1}-T^n}{\tau} + \mathbf{u}\nabla T^{n+1} - \alpha \Delta T^{n+1} = 0$$

Where T^n is T evaluate at time $t_0 + n \cdot dt$. The explicit method is easy to implement but has significant restrictions on temporal stepsize.

On the other hand, the implicit one is more robust but requires solving a linear system for each iteration.



- 1. *Central Finite Difference*, is a second-order method, but is affected by oscillation.
- 2. Upwind scheme, is a first-order method but more stable.

Equations

$$\mathbf{u} \cdot \nabla T = u_1(\partial_x T) + u_2(\partial_y T)$$

1. $\partial_x T = \frac{T_{i+1,j} - T_{i-1,j}}{2 \, dx} + o(||dx||^2)$
2. $\partial_x T = \begin{cases} \frac{T_{i+1,j} - T_{i,j}}{dx} & \text{if } u_1 < 0\\ \frac{T_{i-1,j} - T_{i,j}}{dx} & \text{if } u_1 > 0 \end{cases} + o(||dx||$

Where $T_{ij} = T(x_i, y_j, \bar{t})$ and \bar{t} is fixed. We can use *central differences* when the **Péclet** condition is satisfied: $u_1 dx < 2\alpha$.



We used a **second-order** scheme in space. For the Laplace operator we used classical *Central finite difference*:

$$\alpha \Delta T = \alpha \Big(\frac{T_{i+1,j} - 2T_{i,j} + T_{i-1,j}}{dx^2} + \frac{T_{i,j+1} - 2T_{i,j} + T_{i,j-1}}{dy^2} \Big) \\ + o(||(dx, dy)||^2)$$

for Neumann boundary condition we used one-sided finite difference

$$rac{-3T_{ij}+4T_{ij+1}-T_{ij+2}}{2dy}=g_{\hat{n}}$$

where $g_{\hat{n}}$ is the normal component on the boundary.



Since ν and μ are constant, the Darcy law and the conservation of mass read

$$div(-rac{
u}{\mu}
abla p) = -rac{
u}{\mu}div(
abla p) = 0 \Rightarrow \Delta p = 0$$

the discretization is similar to the Laplacian of *Temperature*. Moreover, we can approximate the velocity vector field \mathbf{u} using the *central finite difference*

$$u_1 = -\frac{\nu}{\mu} \left(\frac{p_{i+1j} - p_{i-1j}}{2dx} \right) \quad u_2 = -\frac{\nu}{\mu} \left(\frac{p_{ij+1} - p_{ij-1}}{2dy} \right)$$

The velocity does not depend on time; it is a stationary flow. We solve these equations only once and then solve the heat equation.

Numerical results: III scenario after 90 days





Numerical results: Evolution





Figure: From left to right: Temperature after 30, 60, 90 days

Numerical results: Animated Evolution



Numerical result: Central F.D. vs Upwind









Conclusions

- $dx \neq dy$
- Upwind vs Central Finite Difference
- spreading of heat fluid
- range of depth

Future works

- $\blacktriangleright \ \mu = \mu(T(\mathbf{x}))$
- $\blacktriangleright \ \nu = \nu(\mathbf{x})$
- II order time (ex. Cranck Nicolson)

References



- A. Coco, J. Gottsmann, F. Whitaker, A. Rust, G. Currenti, A. Jasim, and S. Bunney: "Numerical models for ground deformation and gravity changes during volcanic unrest: simulating the hydrothermal system dynamics of a restless caldera," *Solid Earth, 7, 1–21, 2016.*
- A. Allendes, G. Campana, F. Fuicas, E. Otarola:"Darcy's problem coupled with the heat equation under singular forcing; analysis and discretization".
- Piochi, M., Kilburn, C., Di Vito, M., Mormone, A., Tramelli, A., Troise, C., and De Natale, G.:" The volcanic and geothermally active Campi Flegrei caldera: an integrated multidisciplinary image of its buried structure", *International J. Earth Sci.*, 103, 401–421,2014.
- V. Comincioli, "Analisi numerica : metodi, modelli, applicazioni" Milano [etc.] : McGraw-Hill libri Italia, ©1990.

Thanks for your attention!